

## Lectures 30 & 31

(31.1)

### Reflection and transmission/ refraction of waves at interfaces

In all the phenomena we study at interfaces we imagine that waves of a single frequency  $\omega$  have been incident for a long time so "transient effects" have decayed and the reflected & transmitted waves also have frequency  $\omega$ .

(31.2)

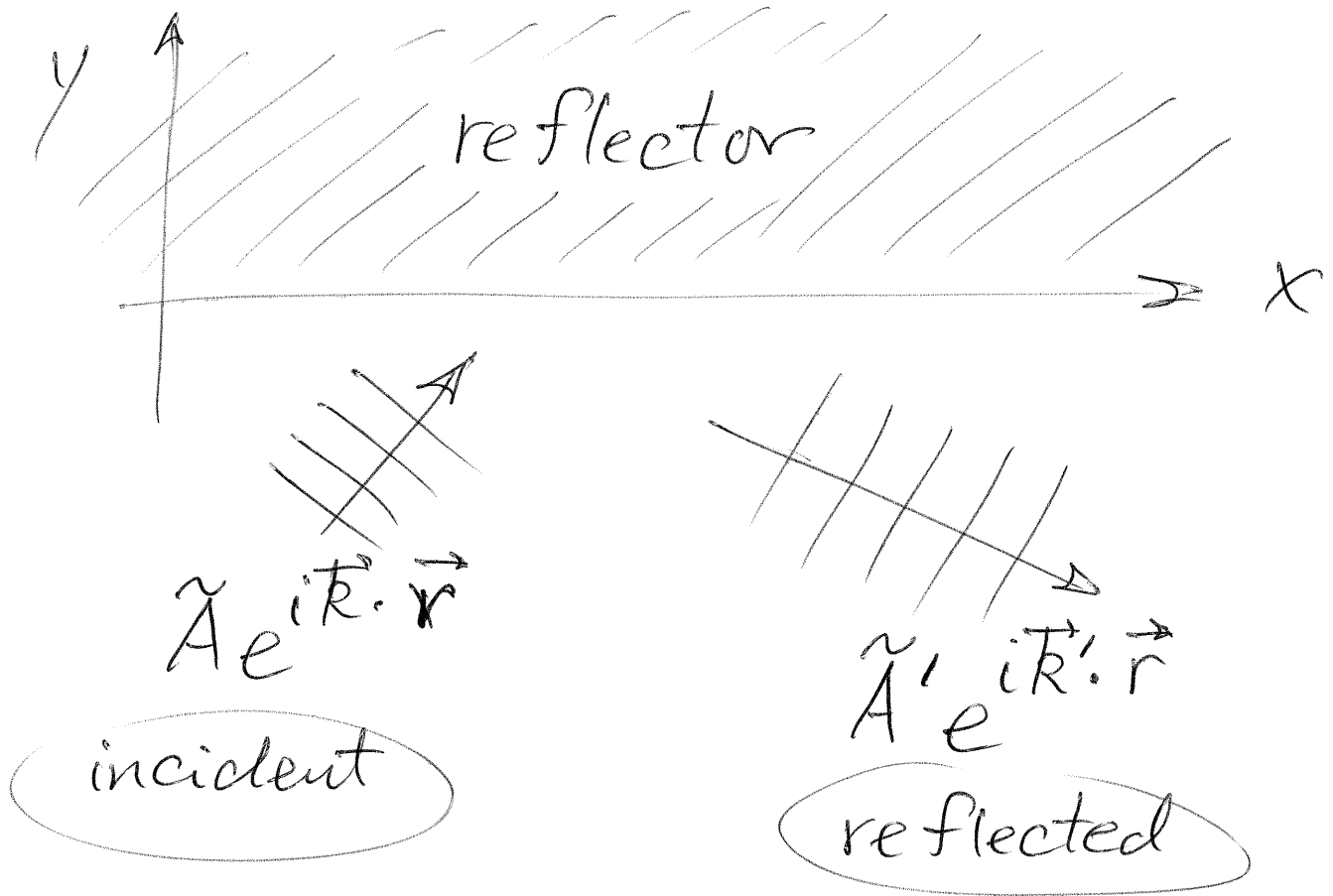
Reflection when the interface forces the amplitude to be zero

This happens in the case of EM waves incident on a good conductor. Recall the property that there can be no static  $\vec{E}$  field inside a conductor. Something close to this is true even when  $\vec{E}$  is oscillating in time (Phys 3327).

Since the time dependence of our waves is always  $e^{-i\omega t}$  we use the compact plane-

wave notation introduced (31.3)

earlier:



• equal freq.  $\Rightarrow |\vec{k}| = |\vec{k}'|$

•  $\tilde{\Psi}(x, 0) = 0 \Rightarrow \Downarrow$

$$\tilde{A} e^{ik_x x} + \tilde{A}' e^{ik'_x x} = 0$$

for all  $x$

The two phasors in (31.4) this equation rotate at different rates with  $x$  if  $k_x \neq k'_x$ , and the equation would not be true for all  $x$  (except in the trivial case  $\tilde{A} = \tilde{A}' = 0$ ). We therefore have  $k_x = k'_x$ . From this we get

$$(\tilde{A} + \tilde{A}') e^{ik_x x} = 0, \text{ all } x$$

and so

$$\tilde{A} + \tilde{A}' = 0.$$

- Since the magnitudes and  $x$ -components of  $\vec{E}$  and  $\vec{E}'$  are equal, their  $y$ -components

must have the same (31.5) magnitude. But only the incident wave may have  $k_y > 0$  (there is only one source at  $y = -\infty$ ), so  $k_y' = -k_y$ .

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In Quantum Mechanics particles are described by waves, and  $\hbar\vec{k}$  is their momentum. The result  $k_x = k_x'$  therefore says that the x-component of momentum is conserved. This is ~~reasonable~~ reasonable considering our interface, which has no structure along x and hence cannot

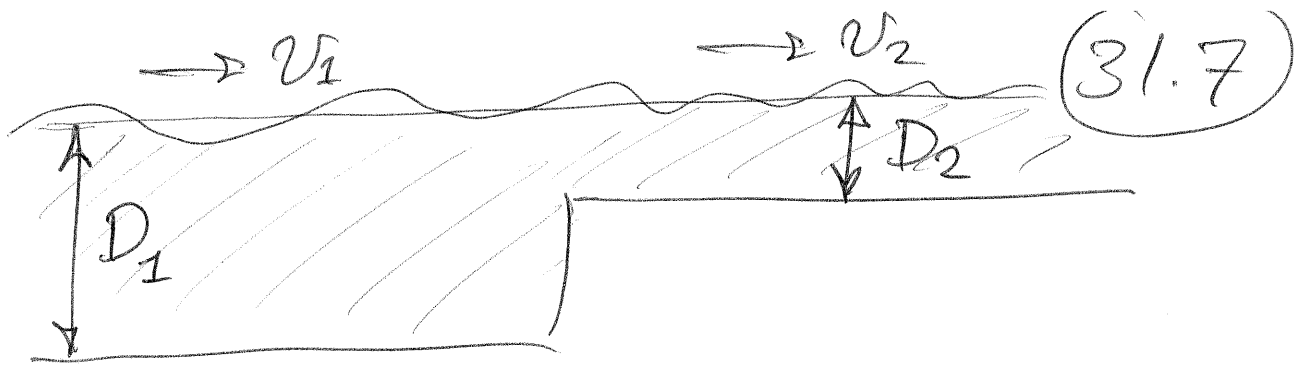
produce forces in  
that direction.

(31.6)

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Normal incidence on interface  
separating media with different  
wave speeds

Imagine surface waves originating far out in the ocean and arriving at a continental "shelf" where the depth is less. Will waves be transmitted into the shallow part? Will there be a reflected wave?



In the limit wavelength much longer than the depth, "shallow" surface waves have this dispersion relation:

$$\omega = \sqrt{gD} k, \quad D = \text{depth}$$

For fixed  $\omega$  in ~~both~~ both depths we have

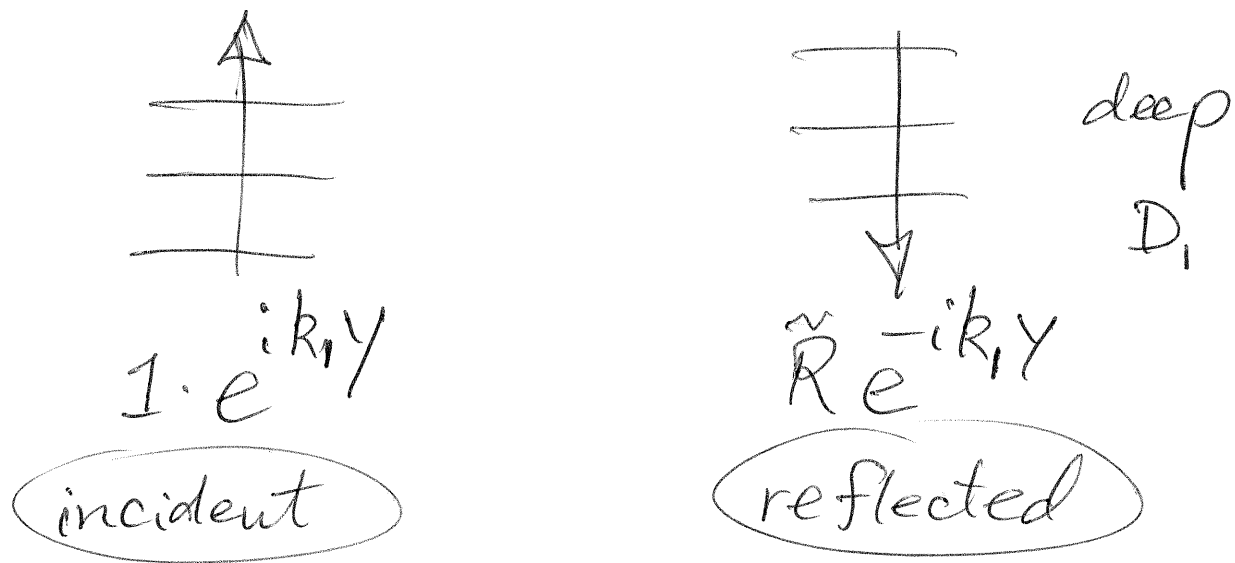
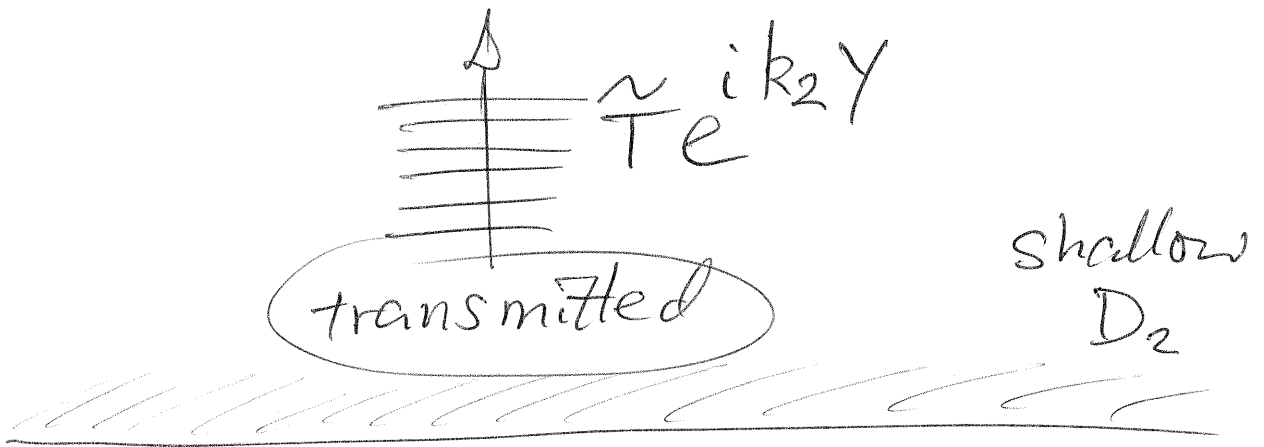
$$\omega = v_1 k_1 = v_2 k_2$$

where  $v_1/v_2 = \sqrt{D_1/D_2}$ .

So waves in the shallow

water move more slowly and have larger  $k$  (shorter wavelength). (31.8)

Top view :



Given  $k_1$  we find  $k_2$  using

$$v_1 k_1 = v_2 k_2 .$$



All that remains is to (31.9)  
determine the amplitudes  
 $\tilde{R}$  &  $\tilde{T}$  (we've set the  
incident amplitude equal to  
1 without loss of generality; for  
any other incident amplitude  
 $\tilde{A}$ , just multiply our results for  
1 as incident:  $\tilde{R} \rightarrow \tilde{A}\tilde{R}$ ,  $\tilde{T} \rightarrow \tilde{A}\tilde{T}$ )

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To determine  $\tilde{R}$  &  $\tilde{T}$  we need  
two equations. These are that  
the wave amplitudes "match" at  
the interface, where "match"  
means

- 1)  $\psi$  is continuous
- 2)  $\frac{d\psi}{dy}$  is continuous

$$1) \Rightarrow 1 + \tilde{R} = \tilde{T}$$

(31.10)

$$(y=0 \text{ limit of } e^{ik_1 y} + \tilde{R} e^{-ik_1 y} = \tilde{T} e^{ik_2 y})$$

$$2) \Rightarrow ik_1 - ik_1 \tilde{R} = ik_2 \tilde{T}$$

$$(y=0 \text{ limit of } \frac{d}{dy}(e^{ik_1 y} + \tilde{R} e^{-ik_1 y}) = \frac{d}{dy}(\tilde{T} e^{ik_2 y}))$$

Solving the two equations in the unknowns  $\tilde{R}, \tilde{T}$ :

$$ik_1(1 - \tilde{R}) = ik_2(1 + \tilde{R})$$

$$i(k_1 - k_2) = i(k_2 + k_1)\tilde{R}$$

$$\Rightarrow \tilde{R} = \frac{k_1 - k_2}{k_1 + k_2}$$

note that  $|\tilde{R}| < 1$ .

$$\tilde{r} = \frac{k_1 + k_2}{k_1 + k_2} + \frac{k_1 - k_2}{k_1 + k_2}$$

$$\tilde{r} = \frac{2k_1}{k_1 + k_2}$$

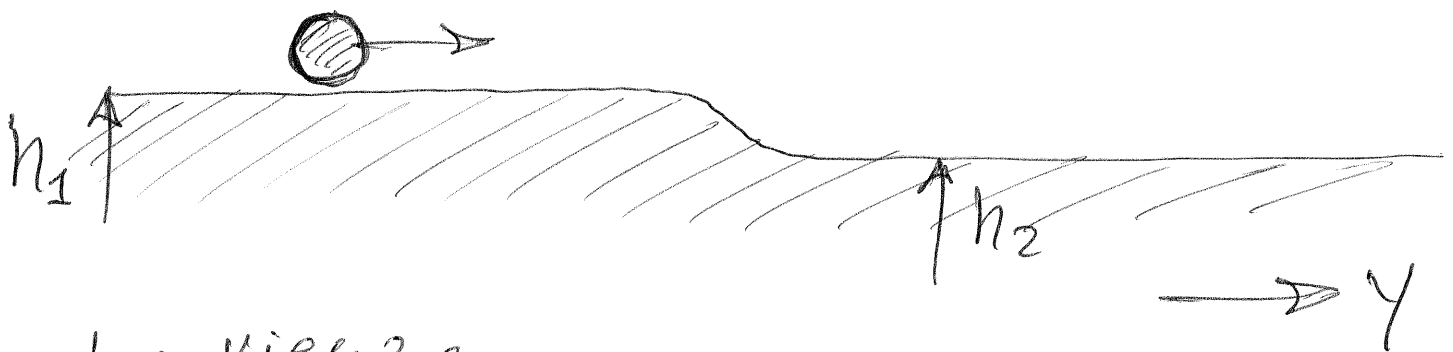
- We get strong reflection, i.e.  $|\tilde{R}| \approx 1$ , if either  $k_1 \gg k_2$  or  $k_1 \ll k_2$ .
- There is poor transmission, i.e.  $\tilde{T} \rightarrow 0$ , if  $k_1 \ll k_2$

Do these results make sense in the context of waves incident from deep/shallow water?

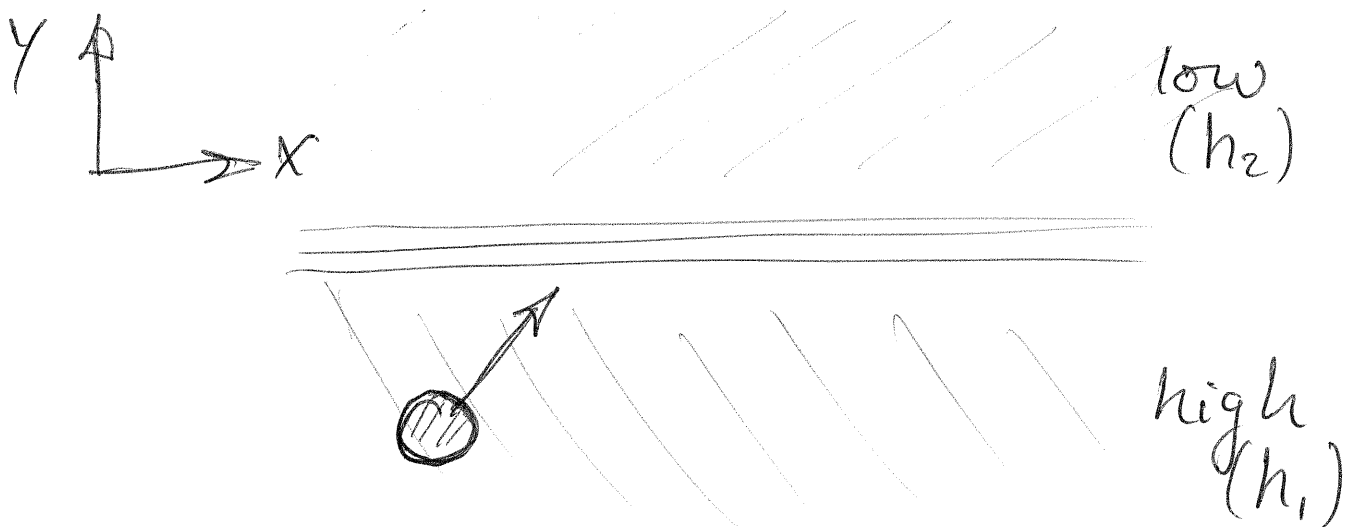
Non-normal incidence:  
refraction

31.12

As a warm-up problem, consider what happens when a marble rolls from one flat surface to another, at a different height:

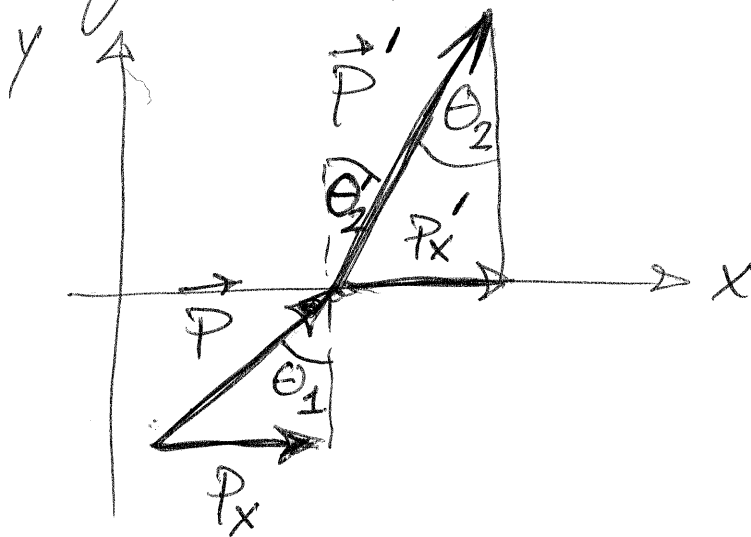


top view:



How do we determine (31.13)  
the marble trajectory in the  
x-y plane using mechanics?

The x-component of the  
marble's momentum is conserved,  
and we can work out what  
happens to the y-component  
using energy conservation:



$$P_x = P'_x$$

or

$$P \sin \theta_1 = P' \sin \theta_2$$

energy conservation:

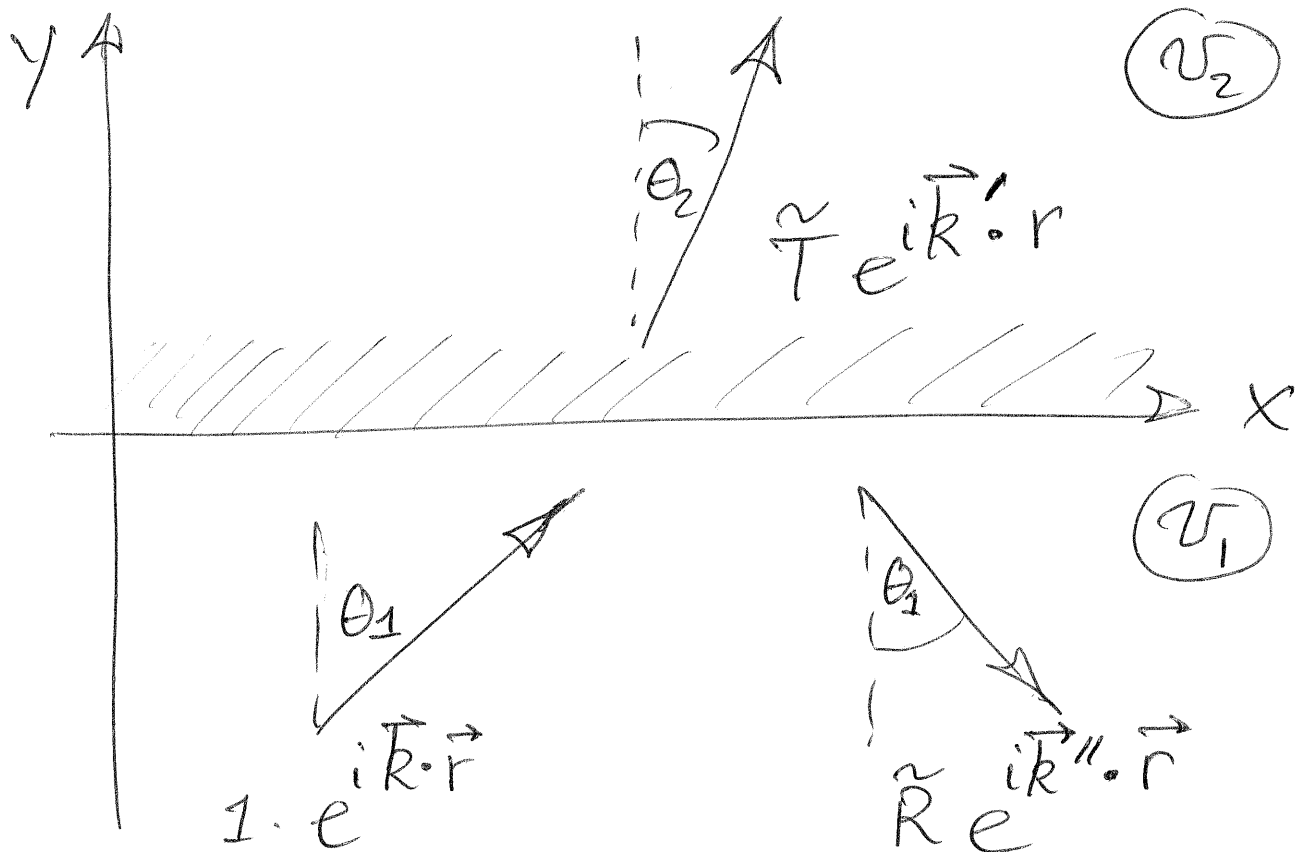
$$\frac{P^2}{2m} + mgh_1 = \frac{P'^2}{2m} + mgh_2$$

$$h_1 > h_2 \Rightarrow P' > P \Rightarrow \theta_2 < \theta_1$$

The situation with waves (3/14) is almost the same. The  $x$ -components of the wave vectors on both sides of the interface are the same and the  $y$ -components are found from the property that the wave has frequency  $\omega$  everywhere ("energy conservation" = "no transient behavior"). There is one difference however: there is a reflected wave as well, and the amplitudes  $\tilde{R}$  &  $\tilde{T}$  of the reflected and transmitted (refracted) waves have to be worked out. But that very thing applies to the marble as well, when it is

described as a quantum wave! (31.15)

We will skip the calculation of the amplitudes  $\tilde{R}$  &  $\tilde{T}$  and focus just on the wave-vectors:



- $k_x = k'_x = k''_x$

- $\underbrace{|\vec{k}| \sin \theta_1}_{k_x} = \underbrace{|\vec{k}'| \sin \theta_2}_{k'_x}$

(31.16)

- $\omega = v_1 |\vec{k}| = v_1 |\vec{k}''| = v_2 |\vec{k}'|$

- $|\vec{k}| = |\vec{k}''|$

- $k_y'' = -k_y$

- $|\vec{k}| \sin \theta_1 = |\vec{k}'| \sin \theta_2$



$$\frac{\omega}{v_1} \sin \theta_1 = \frac{\omega}{v_2} \sin \theta_2$$

In the case of light, the wave ("phase") velocity is written

$$v_1 = \frac{c}{n_1}, \quad v_2 = \frac{c}{n_2}$$

$n_1$  &  $n_2$  = indices of refraction  
in the two media

Hence:  $n_1 \sin \theta_1 = n_2 \sin \theta_2$  (Snell)